

# THE INVESTIGATION OF KIŞLA DOME STRUCTURE IN SOUTHERN PART OF ISPARTA ANGLE. SW TURKEY

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The Isparta Angle is one of the most important geographical structures in SW Anatolia, Turkey developed by bending of Taurides in northern part of Antalya bay. The Edirdir-Kovada graben, which extends through N-S direction and diminishes towards the S, divides Isparta Angle into two equal sections. Mesozoic carbonate sequence forming the western part of Isparta Angle is called "Beydaglan autochthonous". The autochthonous rock sequence forming the western part of Isparta Angle is called "Beydaglan autochthonous". The structural entities within the Isparta Angle extends parallel or semi-parallel to main N-S trending lines. Although a great number of the folds located in the eastern part of Isparta Angle are generally NW-SE extensions, the folds located in the western part of Isparta Angle towards W side until Early Pliocene along N-W Aksu thrust, multi-stage thrusts parallel to Aksu thrust are formed, especially in E side.

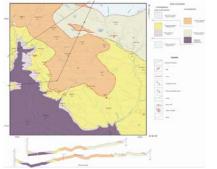
The geological structure known as Kişla dome is located on the N-S extension alkaline Afyon-Antalya volcanic trend. The volcanics consisting of trachyandesites, trachytes, phonolithes and lamproites formed mainly by sub-volcanic dykes, domes and pyroclastic tuffs extend along N-S alkaline volcanic trend within the Isparta Angle. The most important distinguishing formations of the Kişla Dome are circular geometric structures, which are observed by satellite images and sophisticated radial draines systems. The Kişla Dome is formed by the Alsu thrust from E. The grey's pand magnetic data measured on the dome structure range between —20 and —70 mages and 45000 and 47150 fl. respectively. The findings showing that no distinguishable anomalies are present on Kişla dome based on magnetic, gravity and heat flow measurements further indicate that this structure is not formed after magnetic activities and is probably formed after tectoric activities. The E-W thrusts and associated folding systems, which are developed between Middle Econer and Early Miocene rocks on the north of Kişla Dome, show that the region was under the influence of N-S compressions during the Middle to Late Miocene. On the other hand, currently active N-W Alsu thrust

Tauride belt in southern Turkey comprises mainly a stack of imbricated, allochthonous nappes, overlying the autochthonous Tauride carbonate axes. These carbonate axes forms a north pointing cusp named as the Isparta Angle in southwestern Turkey. Ceology and evolution of the Isparta Angle have been investigated by many researchers especially since 1950s (inc. 1977: Nejnoric, 1977: km south of Isparta Irig. 1; tysal 2004; Un the other hand, this obme-like structure take place in the N-5 trending Ayon-Antiaye violenic province of Yagmurlu et al (1997). Atthough having carried out geophysical studies using especially gravity and magnetic methods in the Isparta angle (Ates and Kearey, 1995; Bilm and Ates, 1999; Bilm and Ates, 2004), a dome-like structure was determined first time in this study.



Figure 1. Satellite image of Kışla dome on the Landsat 7-4-1 (1985 August)

The study area is located 20 km south of Isparta and covers an area of 80 km2. Geological map of the study area in 1/10 000 scale, geological cross sections (Fig.2) and generalized tectono-stratigraphic columnar section (Fig.3) were prepared in this study. The rock units are divided into two groups as autochthonous- paraeutochthonous and allochthonous units. Autochthonous-paraeutochthonous units are Devras formation (Early-Late Duressic carbonate rocks). Bydiglian formation (Early-Late Cretaceous carbonate rocks). Yazır limestone (Aquitanian reefal limestone). Burdigalian Aglasun formation including flysch type deposits. Allochthonous units are Antalya nappes and Kaykoy formation fVazur appes in Poisson et al., 2003. Antalya nappes are represented by Isparta Cay formation (Early-Middle Triassic thin bedded cherts and limestones) and ophiolitic complex. Ecoene Kaykoy formation including flysch type deposits rest tectonically on the Aquitanian and Burdigalian units (Fig. 2,3). All the units are in places intruded by trachytic and lamprohyric dykes.



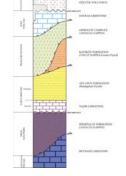
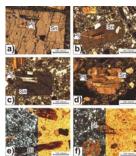


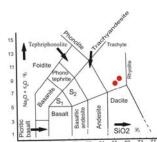
Figure 3. Generalized columnar stratigraphic section of the study area.

# PETROGRAPHY AND GEOCHEMISTRY

Magmatic rocks outcropping around the Kışla dome were investigated in terms of geochemical and petrographic features. Volcanic rocks consist mainly of corroded plagioclase, sanidine, amphibole, pyroxene crystals (Fig. 41. In the SiO<sub>2</sub> vs. Na<sub>2</sub>O + K<sub>2</sub>O (TAS) diagram, volcanic rocks plot in the trachyte area (Fig. 51. Syenitic anclaves, plutonic equivalent of the trachytes, are observed in the volcanic rocks. Therefore, it is thought that syenitic intrusion occurs below the Kışla dome.

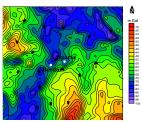


4. In the trachytes outcropping around the Kışla dome a) plagioclase crystals occuring in parallel to margin of sanidine, b.c.d) plagioclase crystals mantled by sanidine, ebictic crystals corroded and crowded with opaque rection products, fl amphibole and sphene crystals corroded along crystal margins. Plj: plagioclase, Sn: sanidine, Bt: biotite, Sf:sphene.



Kişla Domu (Isparta) and surrounding area were investigated for the points of view as geology, tectonics and geophysics. The field studies were carried out between July-August 2007 and numbers of 1761 total magnetic data were collected by Scintrex ENVI MagGrad gradiometer in Kadilar Mahallesi and its vicinity area.

At first, M25 and N25 gravity and aeromagnetic sheets (scaled 1/100.000 and prepared by MTA) were digitized with 1 km data interval. These data were evaluated to determine general tectonic structure of the field by using data processing techiques as filtering, analitical continuations, trend analysis and second derivatives. The gravity data measured on the dome structure range between —80 and —70 mgal (Fig. 6).



The total magnetic field data which were collected in Kadılar Mahallesi were gridded with 25 m. data interval and total magnetic field anomaly map was plotted fig. 7). Gridded magnetic data were evaluated by using data processing techiques as filtering, analitical continuations, trend analysis and second derivatives:

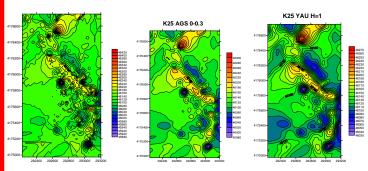


Figure 8. Low pass filtering anomaly map (cut-off frequency 0.0-0.30) (grid interval 25 m., contours in nT) Continuation anomaly map (H=1 grid interval, grid interval 25 m., contours in nT))

As shown Fig.7., the main anomaly structure is trending from SE to NW direction. The same indications are clearly seeing on the low pass filtering and upward analytical continuation anomaly maps (Fig B and Fig 9). The maximum data values are located in Kadilar Mahallesi, north of the magnetic anomaly map. This indicates to existing an uplifiting structure.

In this study, Kışla Domu (Isparta) and surrounding area were investigated for the points of view as geology, tectonics and geophysics. The study area comprises autochthonous-pareautochthonous (Early-Late Jurassic carbonate rocks of Davras formation, Aquitanian refelf limestone of Varz limestone, Burdigalian Aglasun formation including flysch type deposits) and allochthonous units (Antalya nappes and Kayıkoy formation) are observed in the study area. All the units are in places intruded by tractlycic and lamprohyric dyles. The gravity and magnetic data measured on the dome structure range between —50 and—70 mgal and 46000 and 47150 fl. respectively.

The findings showing that no distinguishable anomalies are present on Kışla dome based on magnetic, gravity and heat flow measurements further indicate that this structure is not formed after magmatic activities and is probably formed after tectonic activities.

## **ACKNOWLEDGEMENTS**

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